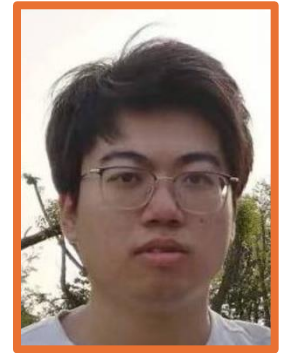


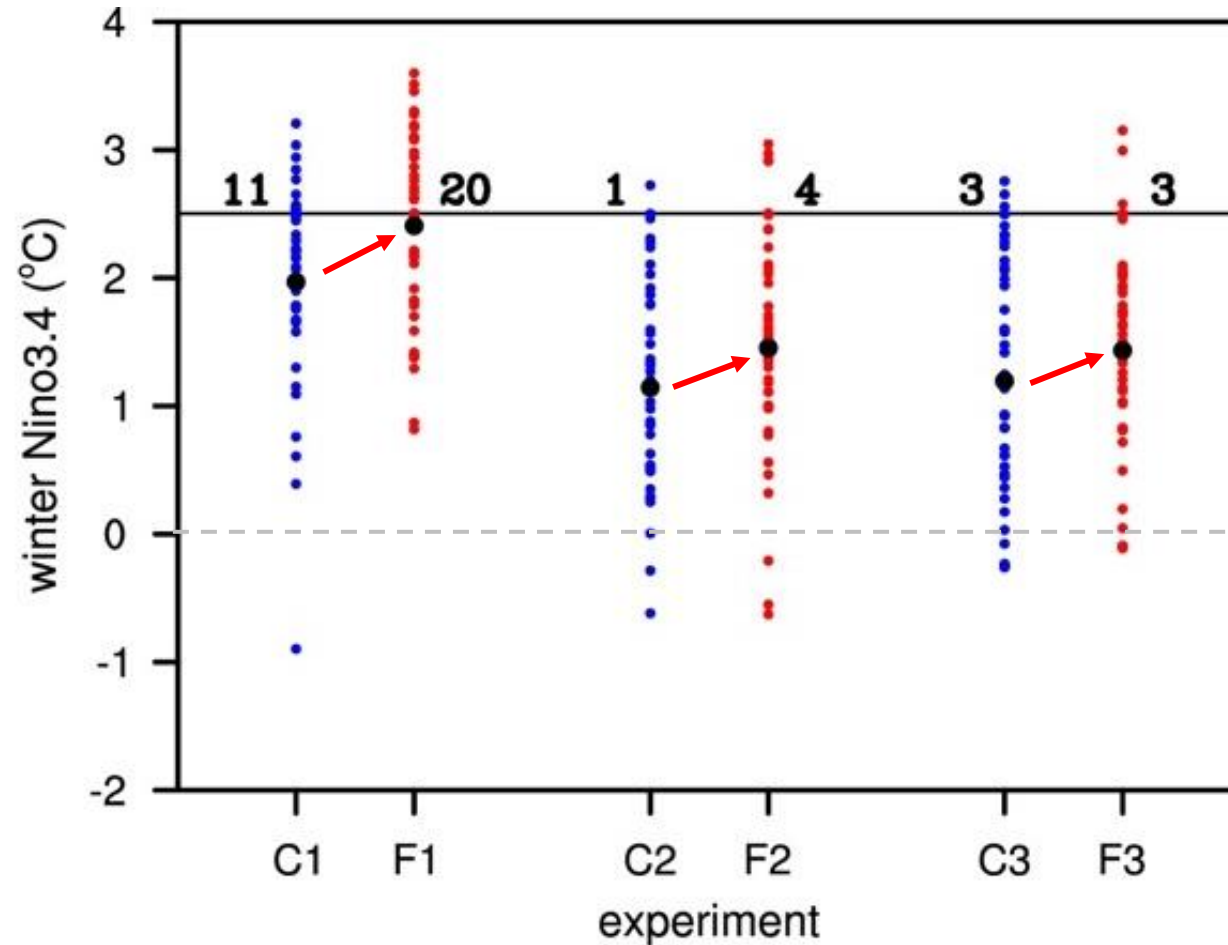
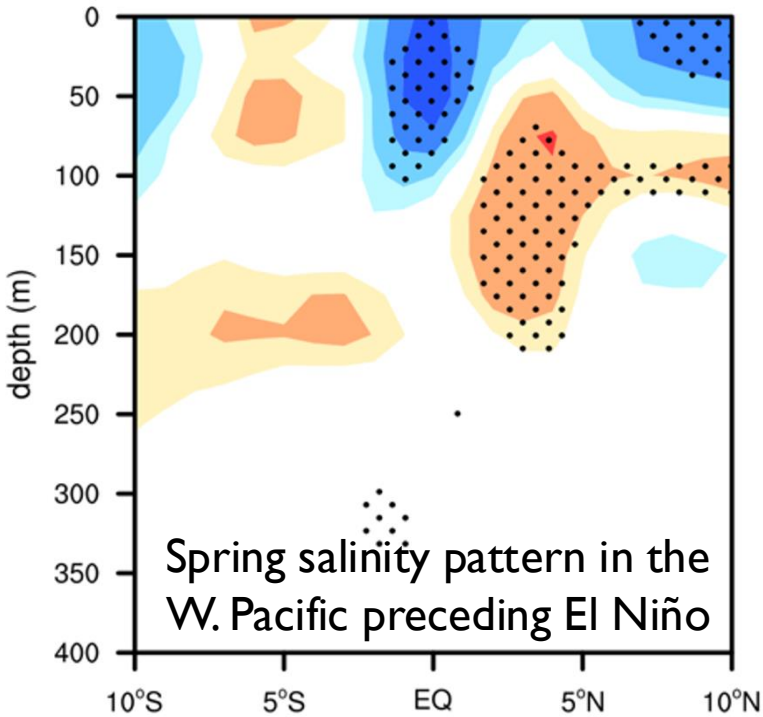
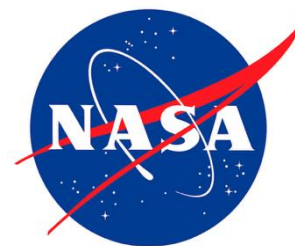
# Salinity-Induced Eastward Flow in Boreal Spring Favors Extreme El Niño

Shineng Hu<sup>1</sup>, Shizuo Liu<sup>1</sup>, Michael McPhaden<sup>2</sup>

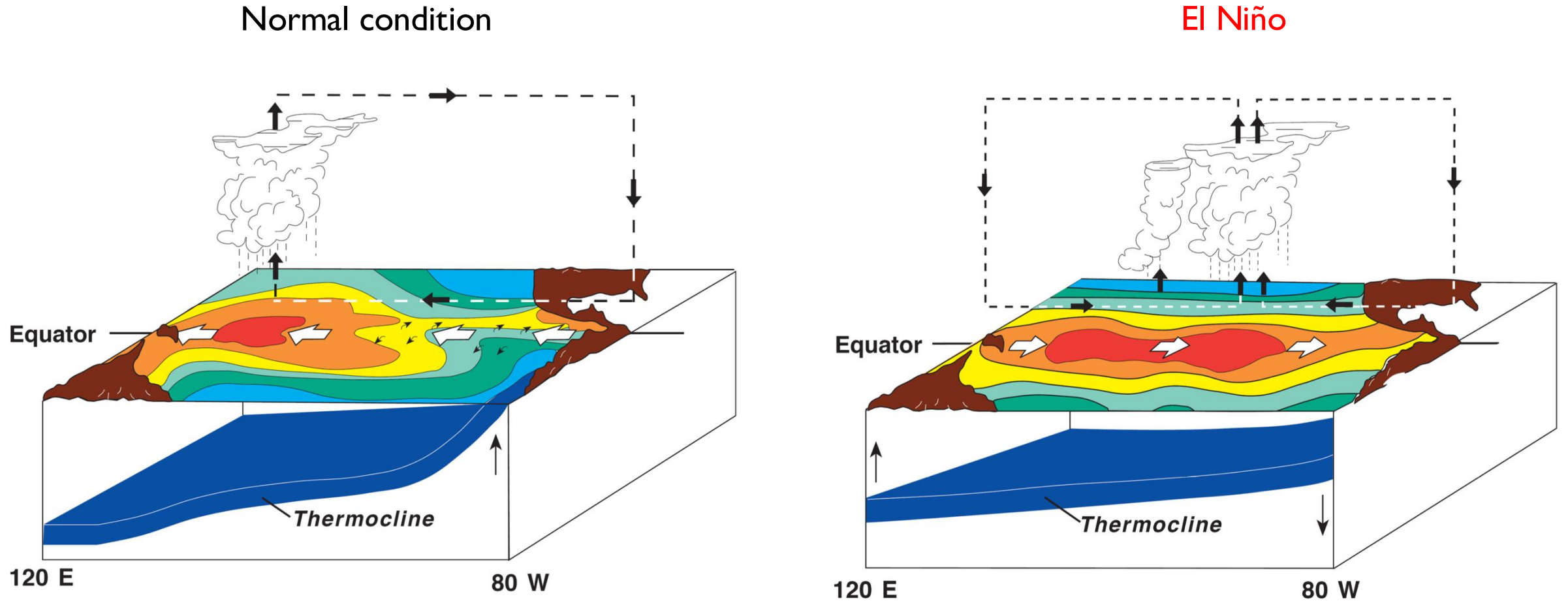
<sup>1</sup>Duke, <sup>2</sup>NOAA/PMEL



Dr. Shizuo Liu



# El Niño-Southern Oscillation (ENSO)

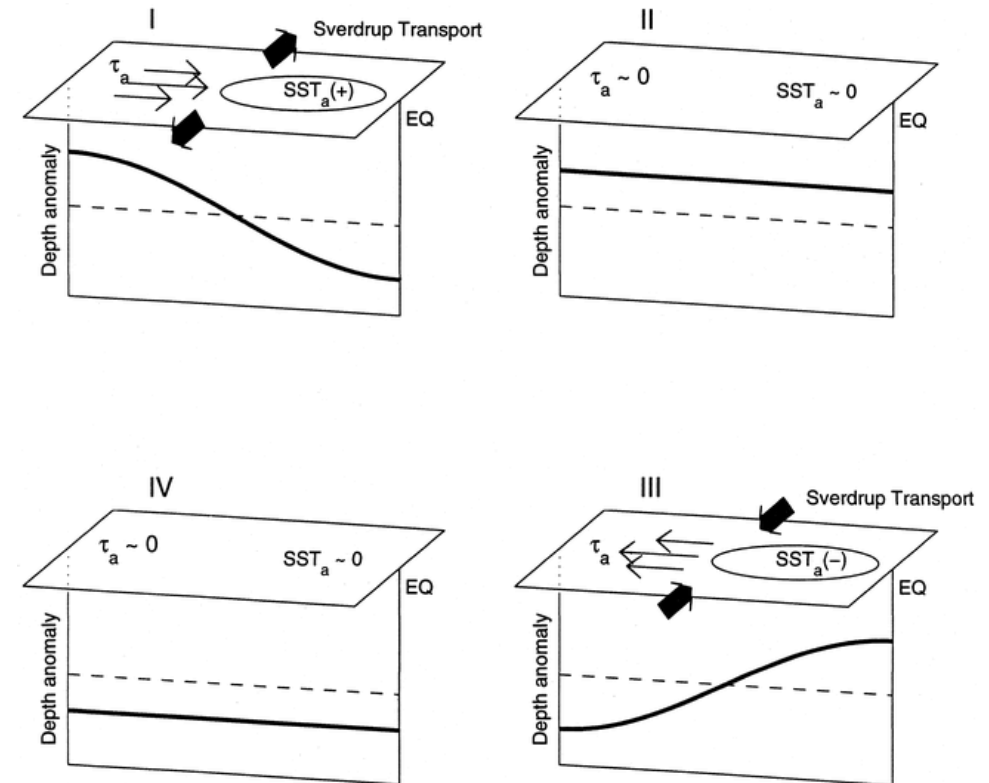


McPhaden et al. (1998)

## A recipe of ENSO:

- 1) Bjerknes feedback: a positive feedback involving sea surface temperature, wind, and thermocline  
(Bjerknes 1969)  
*(How does El Niño/La Niña amplify?)*
- 2) “Recharge Oscillator”: a delayed, negative feedback involving subsurface ocean heat change  
(Jin 1997; Meinen and McPhaden 2000; Vialard et al. 2025)  
*(How does El Niño/La Niña change phase?)*
- 3) Westerly wind bursts: a trigger of El Niño events & a part of the Bjerknes feedback as a multiplicative noise  
(e.g. Kessler et al. 1995, Yu et al. 2003; Lengaigne et al. 2004; Eisenman et al. 2005; Hu et al. 2014; Hu and Fedorov 2019)  
*(What maintains and energizes the oscillation?)*

## Schematic of “Recharge Oscillator”

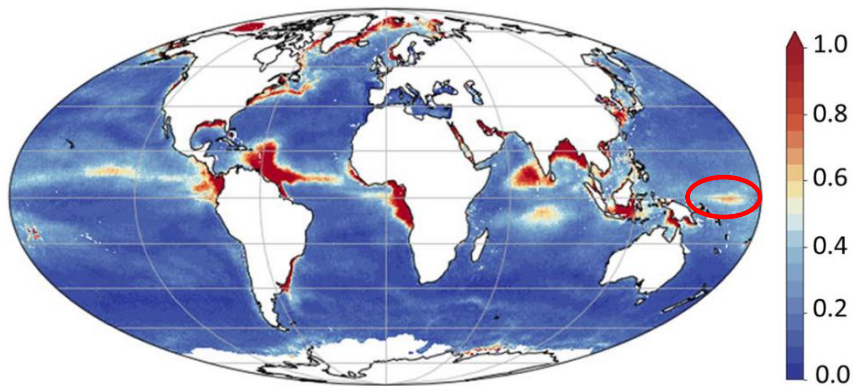


Jin (1997); Meinen and McPhaden (2000)

**What is the role of ocean salinity in ENSO dynamics?**

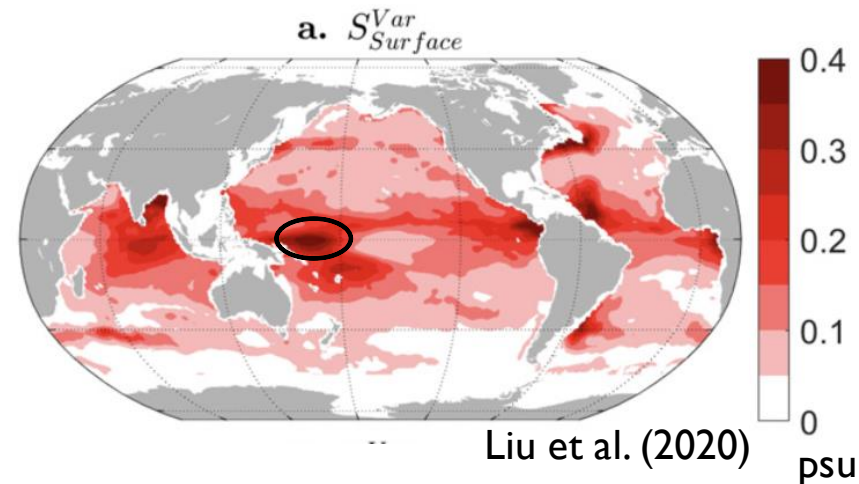
# A role of ocean salinity in ENSO dynamics?

Satellite-based surface salinity variability



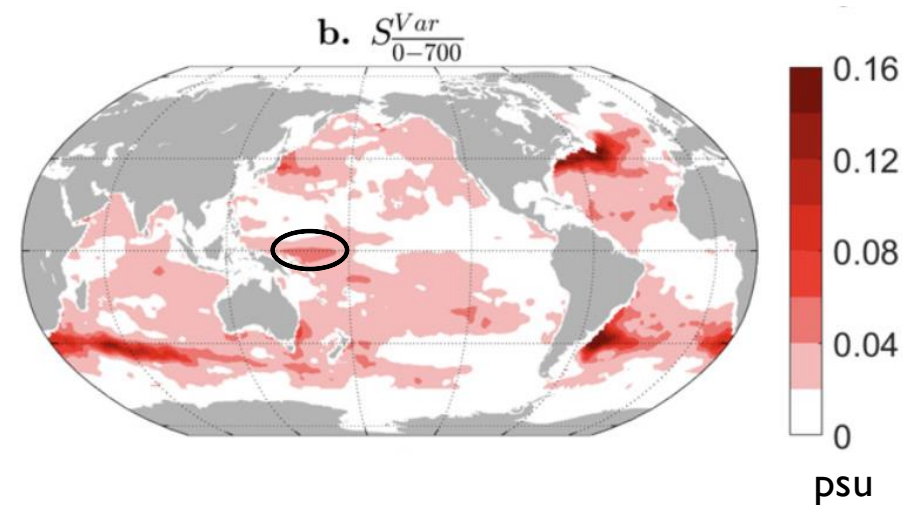
Boutin et al. (2021)

Argo-based surface salinity variability

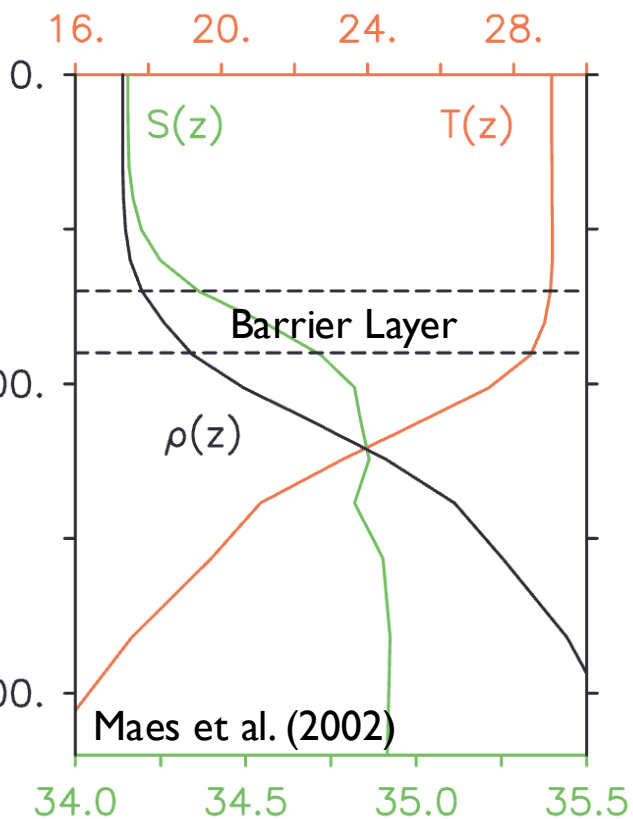


Liu et al. (2020)

Argo-based upper ocean salinity variability



Can ocean salinity **variability** influence ENSO dynamics?

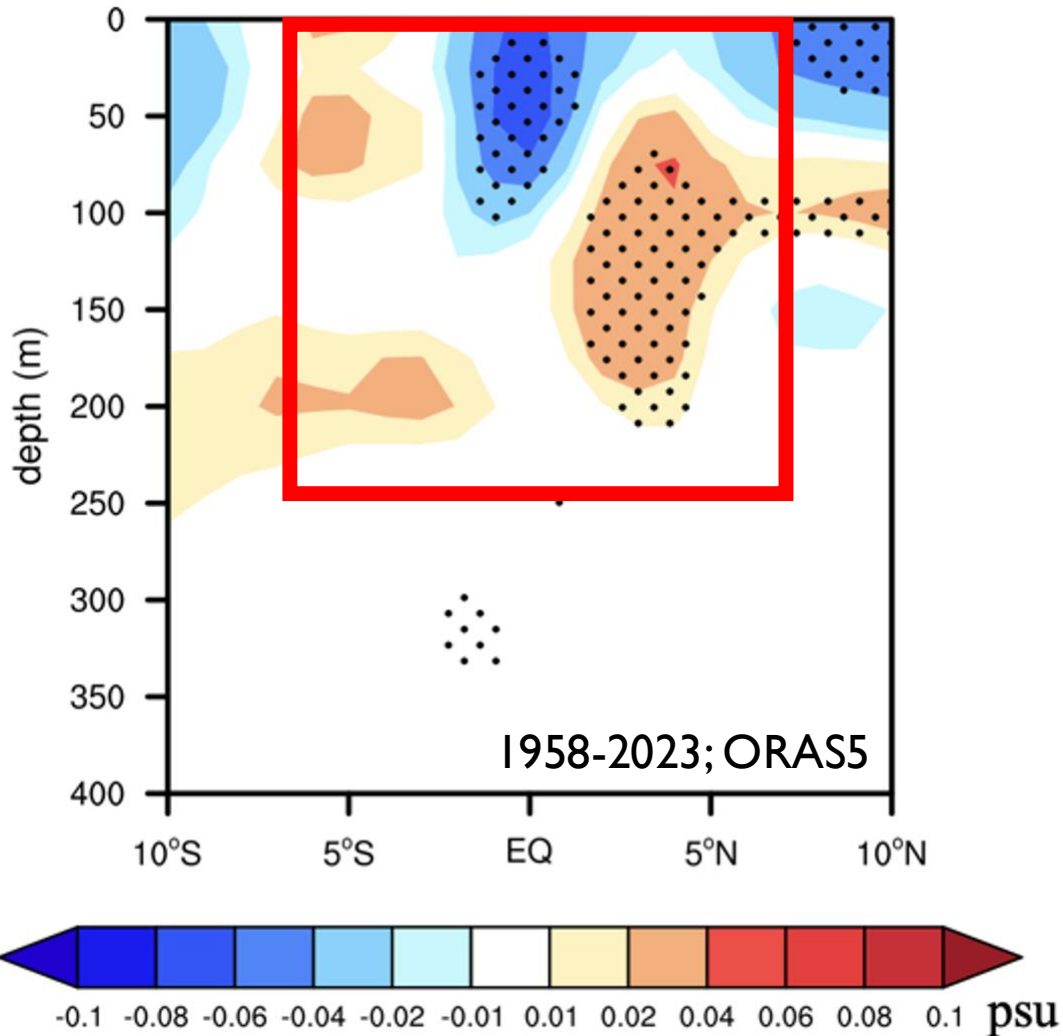


Maes et al. (2002)

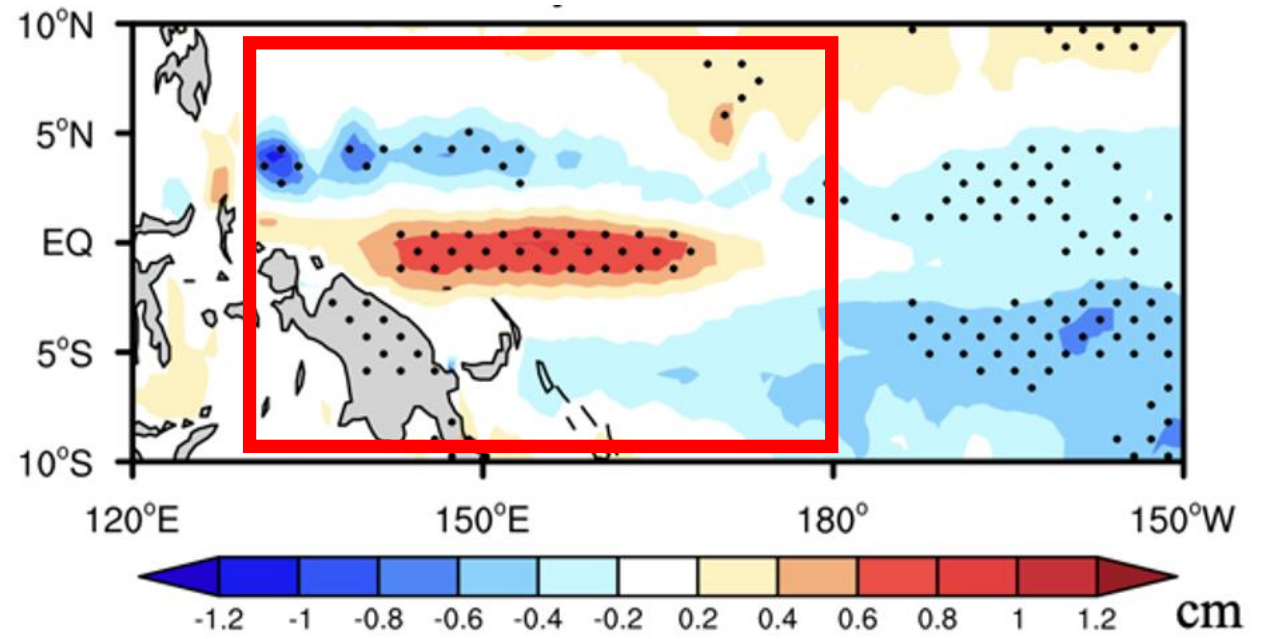
- Ocean barrier layer in the western Pacific hinders entrainment of subsurface cool water and preserves ocean heat content. (e.g., Vialard and Delecluse 1998; Maes et al. 2002;)

# March-May salinity pattern preceding El Niño events (Regression onto normalized December-February Niño3.4)

Salinity anomalies in the W. Pacific (130°E-180°)



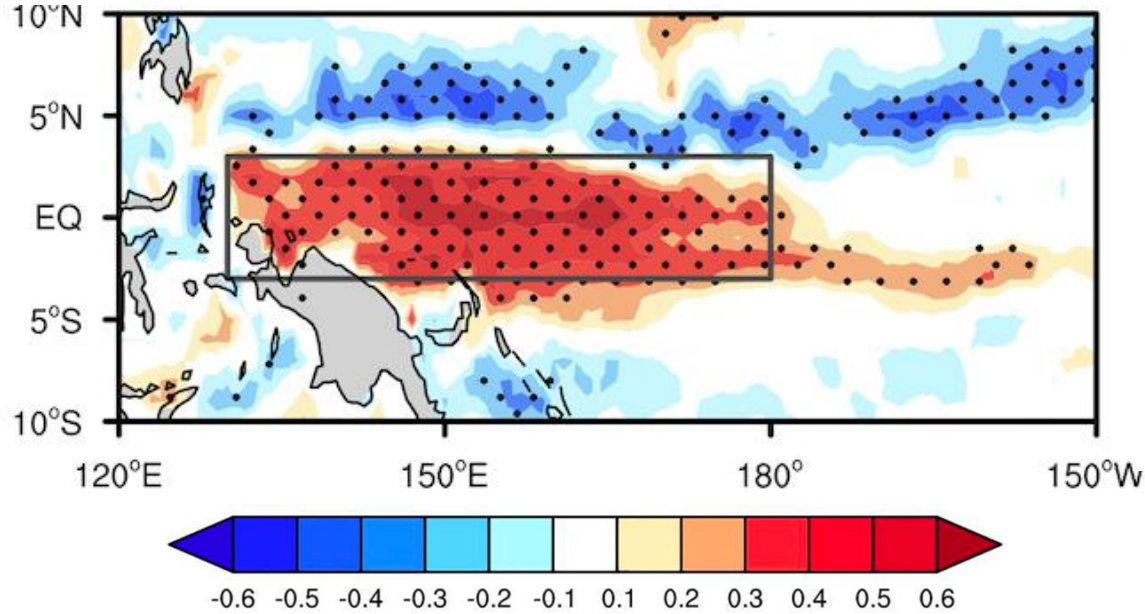
Steric sea level due to subsurface salinity anomalies (0-400m)



$$\xi_S = - \int_{z=0}^H \frac{\rho(\bar{T}, S, p) - \bar{\rho}(T, S, p)}{\bar{\rho}(T, S, p)} dz$$

# Surface zonal (semi)geostrophic current associated with the spring salinity pattern

Correlation between spring (MAM) salinity-induced surface zonal geostrophic current and winter (DJF) Niño3.4

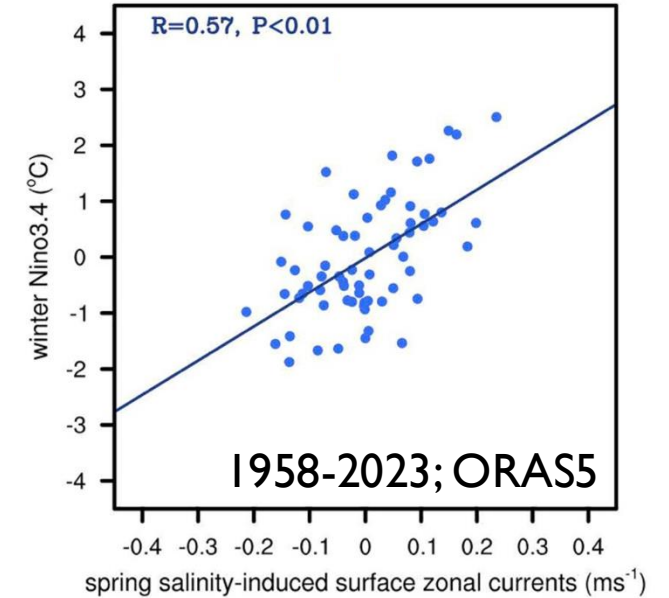


Off the equator:

$$u_g = -\frac{g}{f} \frac{\partial \xi}{\partial y}$$

At the equator:

$$u_{sg} = -\frac{g}{\beta} \frac{\partial^2 \xi}{\partial y^2}$$

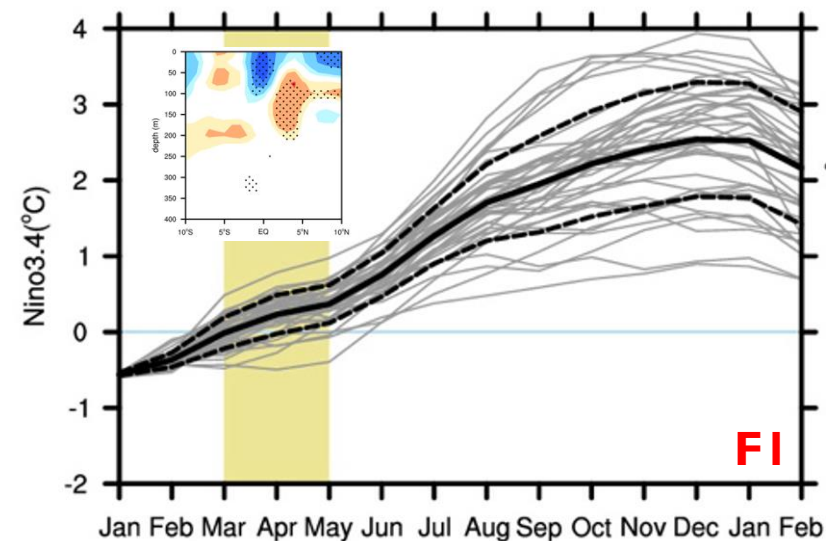
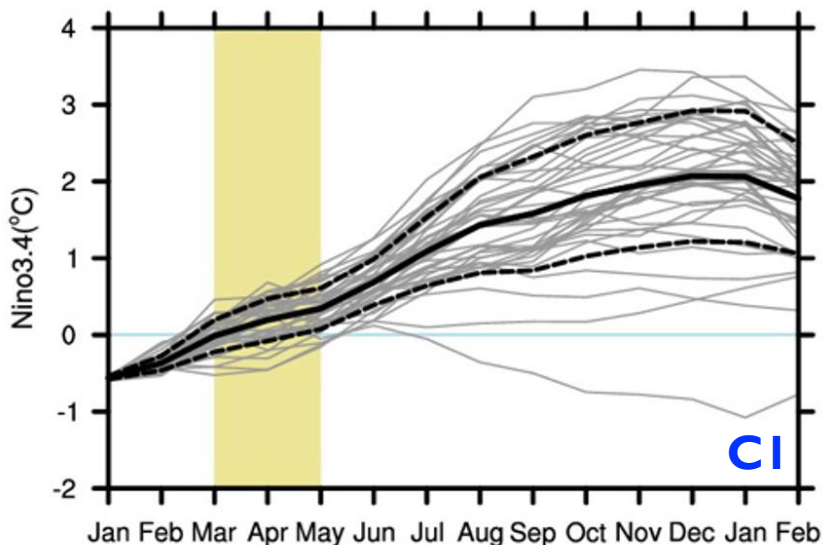


**Hypothesis:**

**The observed spring salinity pattern can facilitate El Niño development through dynamical processes (geostrophic currents, Kelvin waves, etc.)**

# Ocean-atmosphere coupled simulations with imposed salinity anomalies

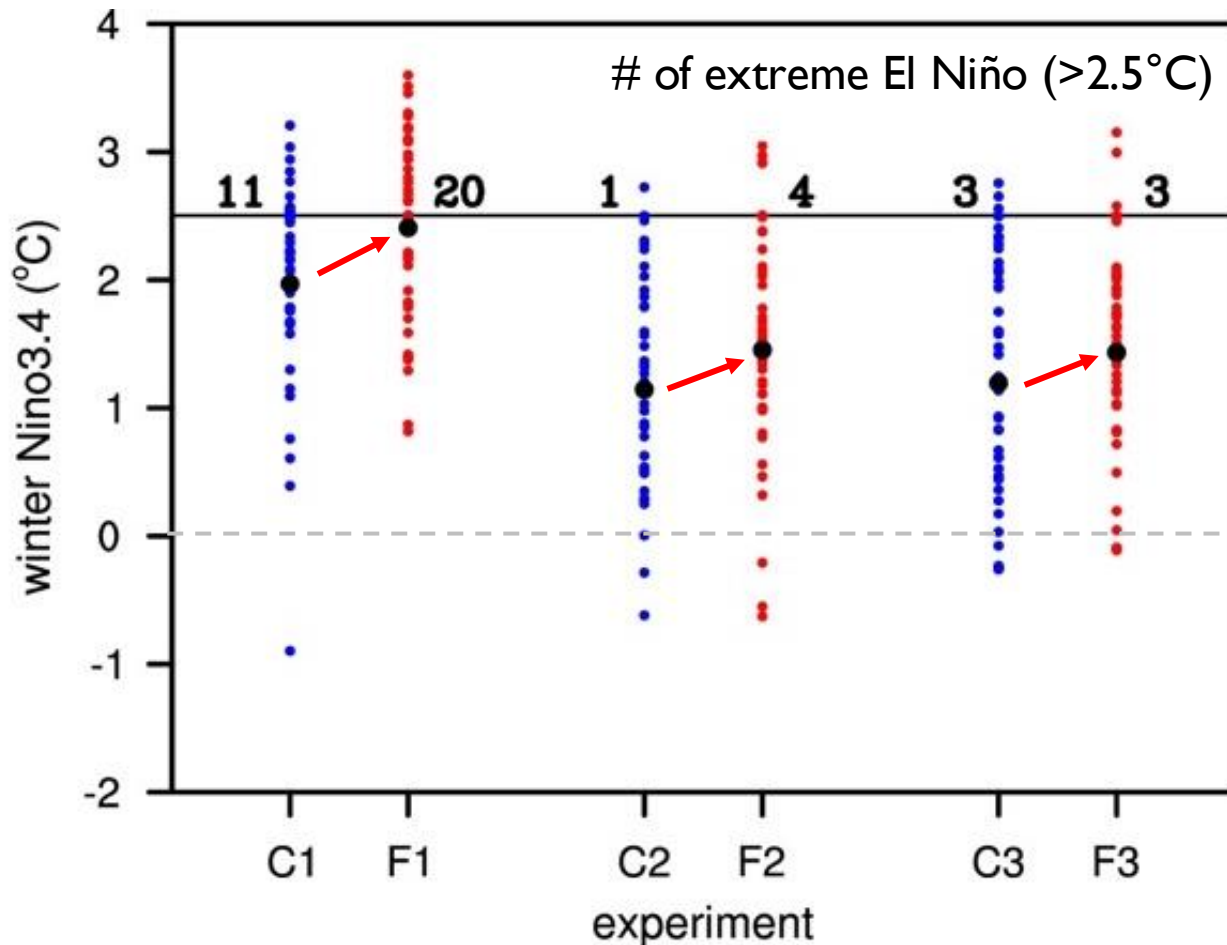
- The Community Earth System Model version 2.1 (CESM2.1)
- Step 1: From a long pre-industrial run, pick a “recharged” ocean initial condition favorable for El Niño development
- Step 2: **Control Ensemble (CI)** & **Salinity-Forced Ensemble (FI)**:
  - Same ocean initial condition (“recharged”)
  - **FI**: salinity forcing applied from **March to May** in the **tropical western Pacific** ( $10^{\circ}\text{S}$ – $10^{\circ}\text{N}$ ,  $130^{\circ}\text{E}$ – $180^{\circ}$ ) between **0-400 m** with a magnitude equivalent to **two times of the regression** of the salinity against the subsequent winter Niño3.4
  - **CI** & **FI**: 40 ensemble members each (slight perturbations of atmospheric initial state)
  - Each simulation: 14 months (from Jan 1 to Feb 28 of the following year).



# Ocean-atmosphere coupled simulations with imposed salinity anomalies

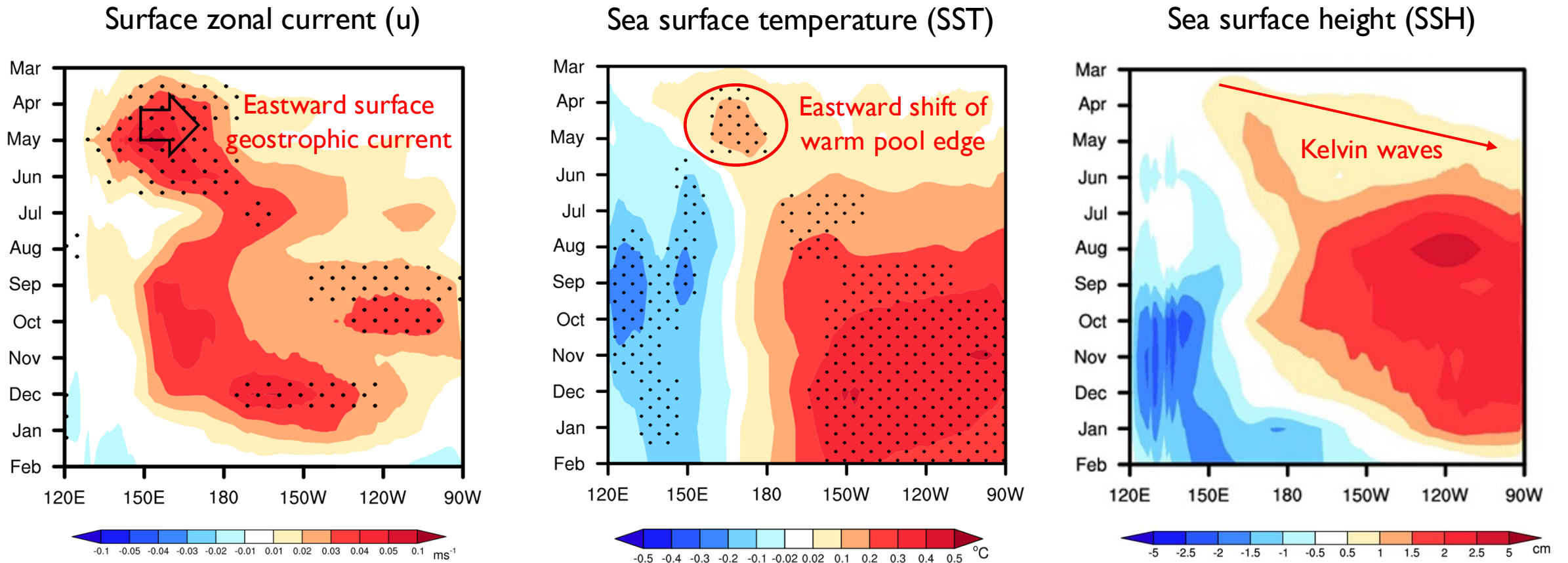
- The Community Earth System Model version 2.1 (CESM2.1)
- Step 1: From a long pre-industrial run, pick a “recharged” ocean initial condition favorable for El Niño development
- Step 2: **Control Ensemble (C1)** & **Salinity-Forced Ensemble (F1)**:
  - Same ocean initial condition (“recharged”)
  - **F1**: salinity forcing applied from **March to May** in the **tropical western Pacific** ( $10^{\circ}\text{S}$ – $10^{\circ}\text{N}$ ,  $130^{\circ}\text{E}$ – $180^{\circ}$ ) between **0-400 m** with a magnitude equivalent to **two times of the regression** of the salinity against the subsequent winter Niño3.4
  - **C1** & **F1**: 40 ensemble members each (slight perturbations of atmospheric initial state)
  - Each simulation: 14 months (from Jan 1 to Feb 28 of the following year).
- Step 3: Repeat Steps 1 & 2 to generate **(C2, F2)**, **(C3, F3)**.

# Winter (DJF) Niño3.4 response to spring salinity forcing (forcing = salinity pattern for a 2-std El Niño)



- On average, wintertime Niño3.4 increases by +0.32°C (from +1.44°C to +1.76°C).
- The probability of extreme El Niño (wintertime Niño3.4 > +2.5°C) almost doubles, from 15 to 27 out of 120.

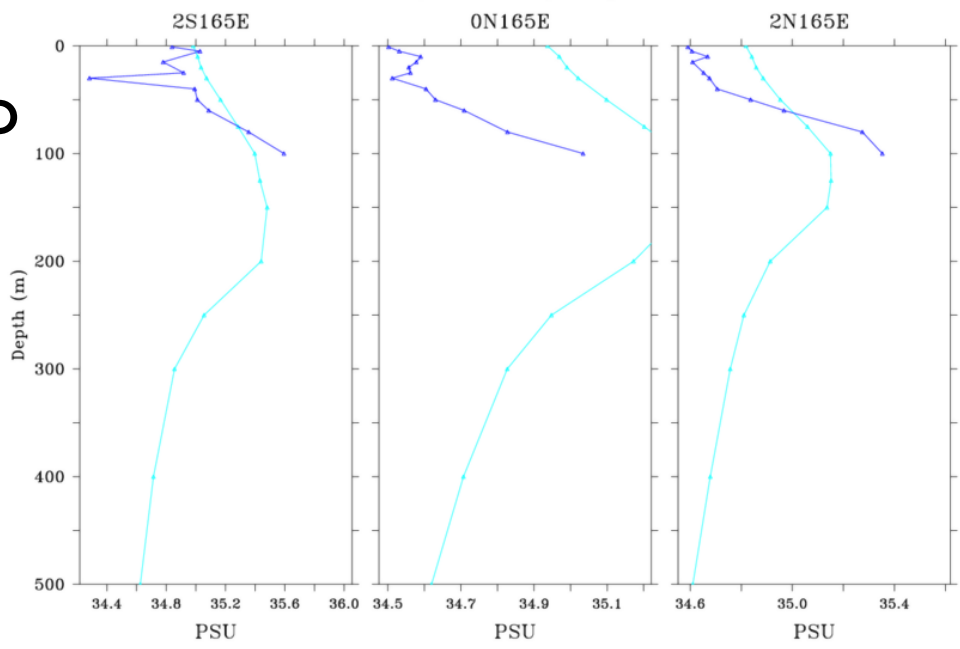
# Hovmoller diagrams (ensemble mean)



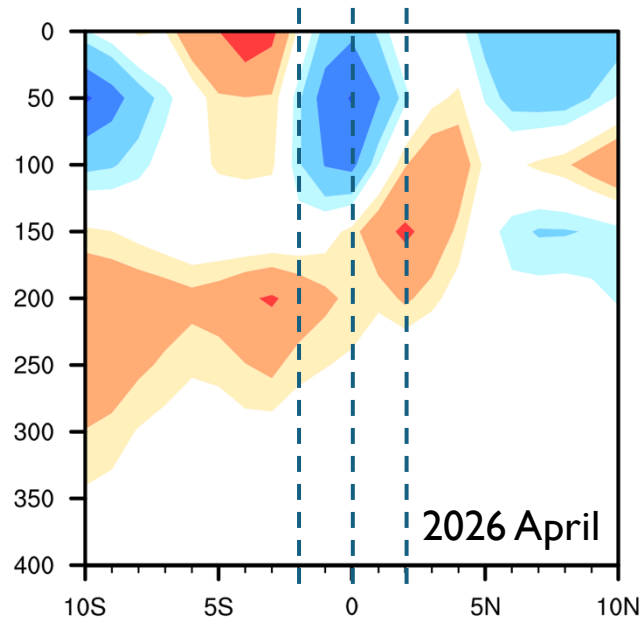
- Imposed salinity anomalies  $\rightarrow$  Eastward surface geostrophic currents & higher SSH in the W. Eq. Pacific.
- Eastward flow  $\rightarrow$  Shifted warm pool edge  $\rightarrow$  Warm SST anomalies near the Dateline
- Higher SSH  $\rightarrow$  Downwelling Kelvin waves  $\rightarrow$  Warm SST anomalies in the E. Eq. Pacific
- Warm SST anomalies trigger positive Bjerknes feedback

# 2026 El Niño

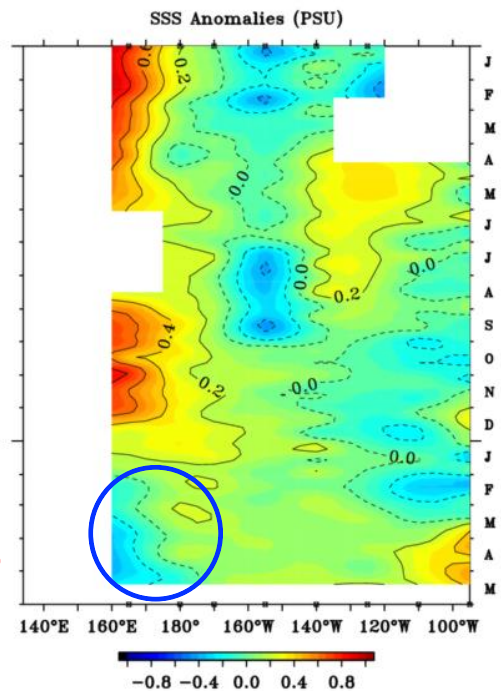
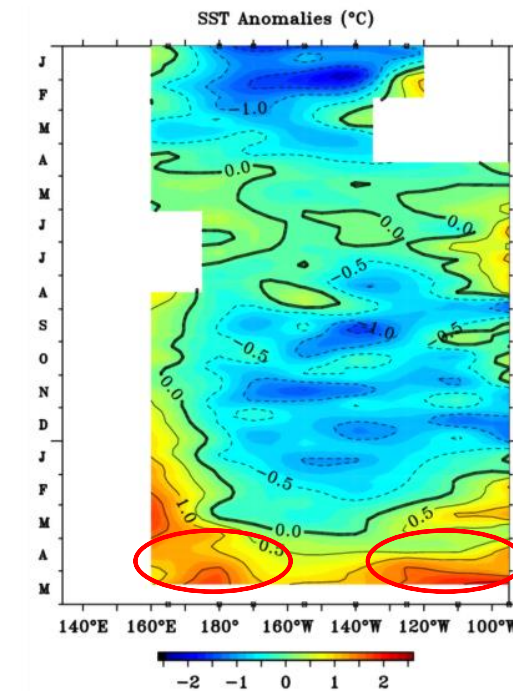
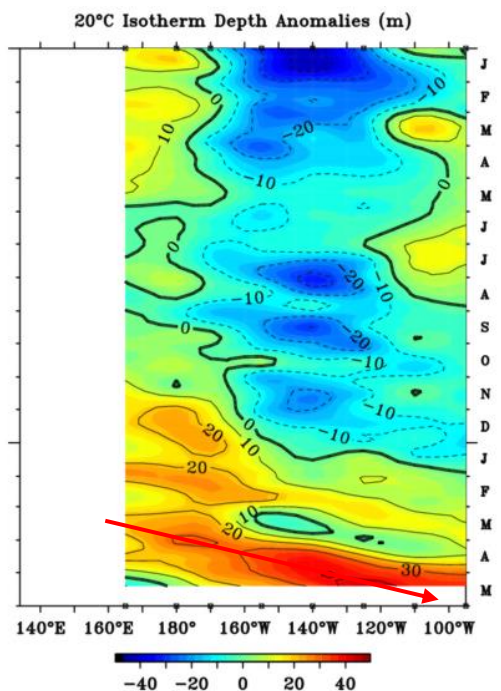
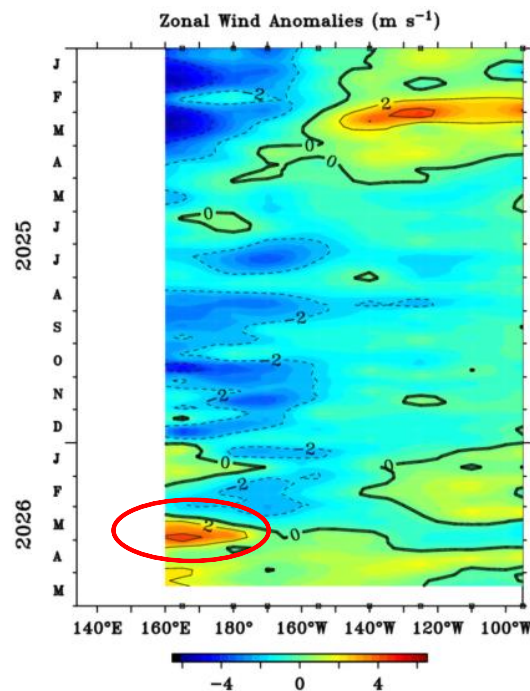
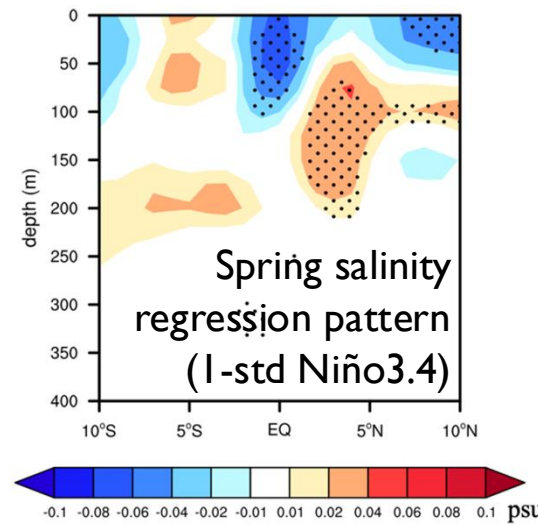
Monthly Salinity April 2026



Global Tropical Moored Buoy Array Program Office, NOAA/PMEL

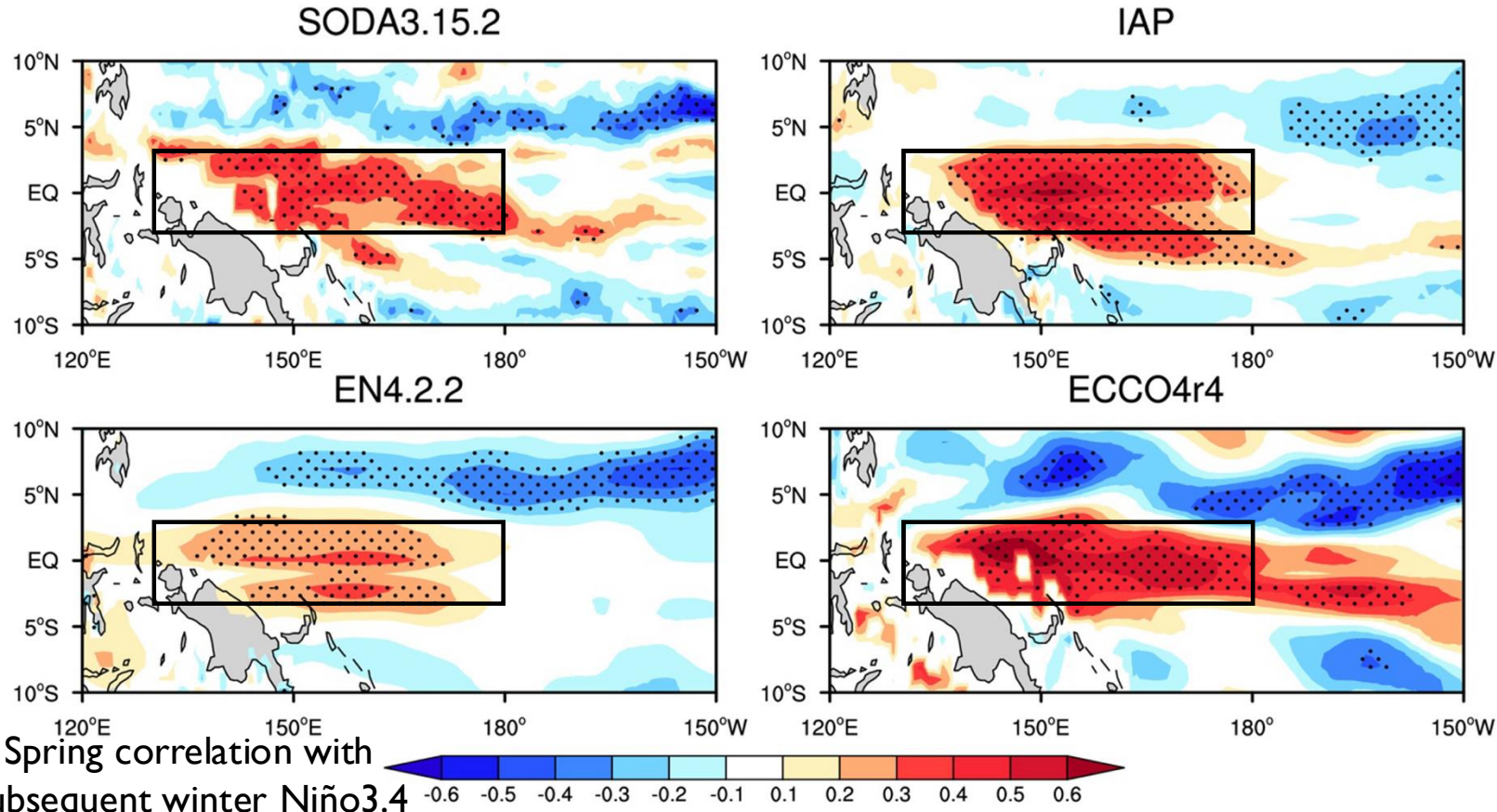


May 16 2026

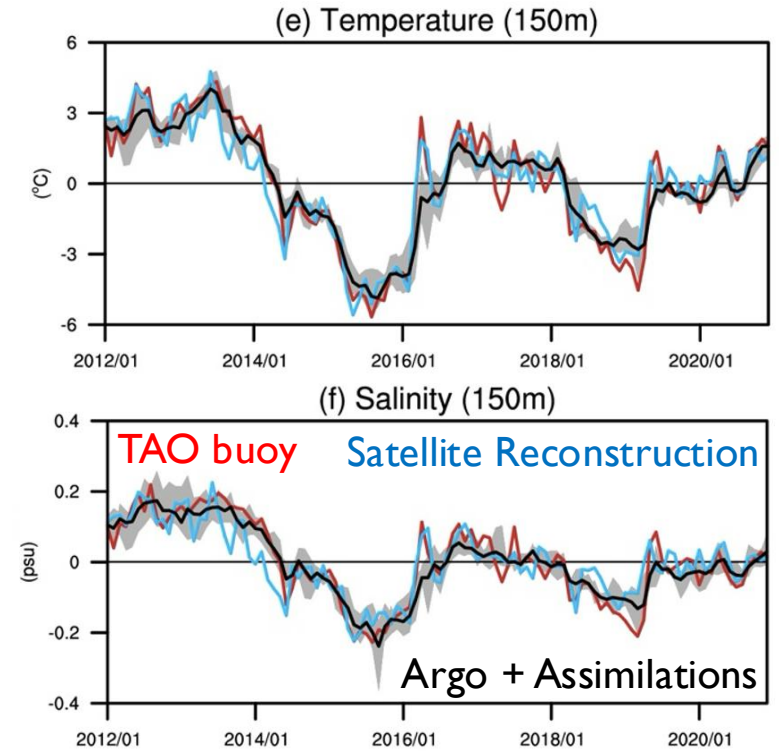


# Satellite-based subsurface T & S data are needed

Diverging signals of ENSO-related salinity-induced surface currents among observational products



Reconstructed T(z) & S(z) in the upper ocean (0-400 m) based on satellite measurements of SSH, SST, and SSS (Liu and Hu, 202X)

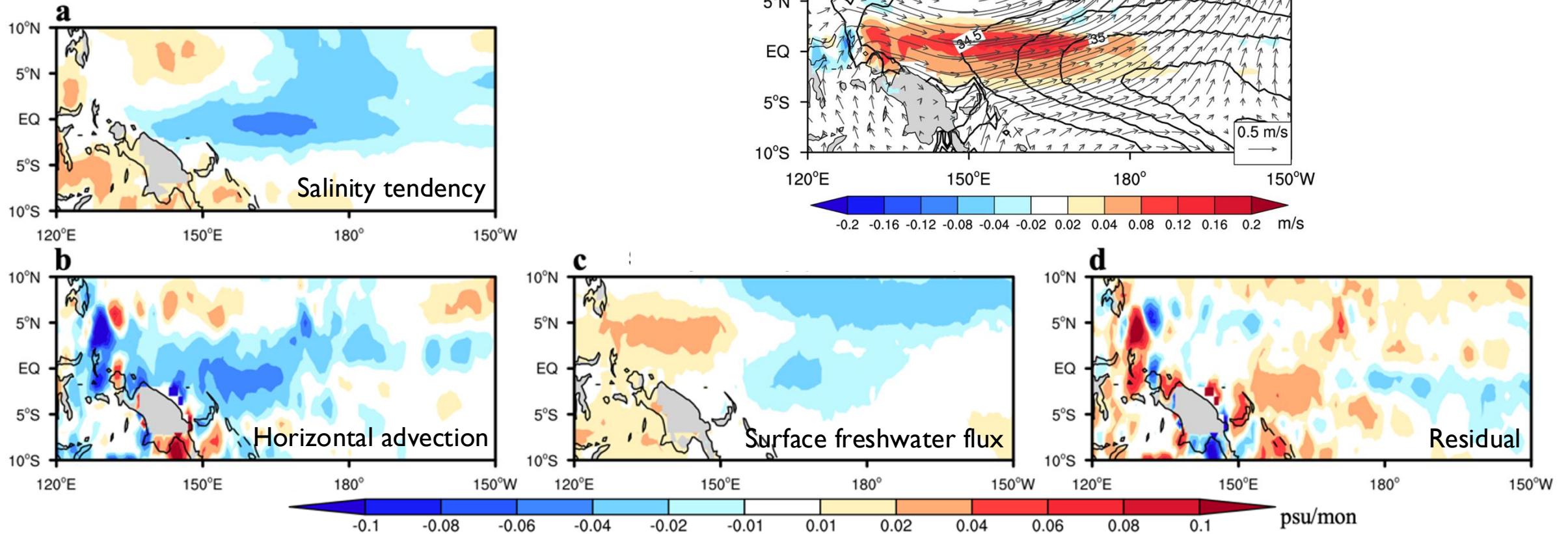


An example in W. Pacific (8°N, 137°E) at 150 m

# Summary

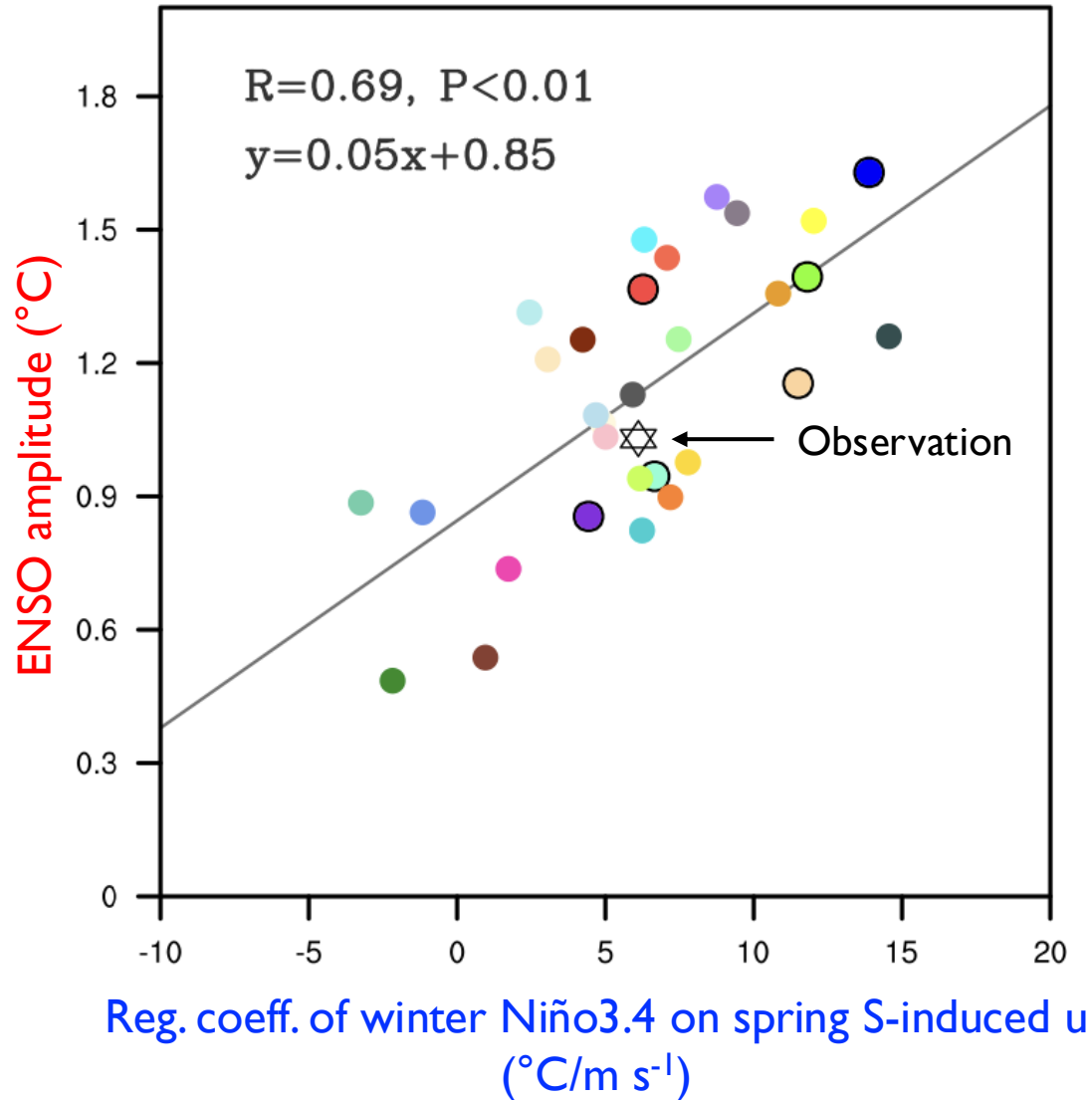
- A robust **spring salinity pattern** with fresher equatorial and saltier off-equatorial anomalies in the western Pacific **precedes El Niño events**.
- These boreal spring salinity anomalies can influence El Niño development through salinity-induced **surface zonal geostrophic currents**.
- Large-ensemble simulations (w/ CESM2) indicate this salinity mechanism can increase El Niño intensity by  $\sim 20\%$ .
- This spring salinity pattern has been identified in March-April 2026, which may contribute to the increase of winter Niño3.4 by  $\sim 0.8^\circ\text{C}$  (CESM2-based rough estimate).

# Spring mixed layer salinity budget (Regression onto normalized December-February Niño3.4)



- The spring salinity pattern preceding an El Niño results mainly from horizontal advection driven by westerly wind bursts and associated surface freshwater flux. (Therefore, probably a booster rather than an independent precursor.)

# CMIP6 model spread in salinity-ENSO connection



- A model that has a **higher sensitivity** of the winter Niño3.4 index to the preceding spring western tropical Pacific salinity-induced surface zonal currents tends to produce a **stronger ENSO**.

